

# On Geoengineering and Implementing an Albedo Solution with Urban Heat Islands Global Warming and Cooling Estimates

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## Abstract

Solar geoengineering is vital in Global Warming (GW) as results can reverse trends and reduce the probability of a tipping point. This paper focuses on geoengineering and implementation of a surface solar geoengineering solution to global warming. Although an albedo solution is reasonably practical, work in this area appears stagnant and even implementing Urban Heat Island (UHI) cool roofs on a global level has not yet been widely adopted. This paper provides basic modeling and motivation by illustrating the potential impact for reverse forcing. We provide insights into “Earthly components” that can be utilized to increase the opportunity for reducing climate change. Modeling shows that by solar geoengineering hotspots with large heat capacities, such as UHIs, and mountain region, the effective area could be roughly 11 times smaller than nominal non-hotspot regions in influencing global warming. We find that between 0.2 and 0.5% of the Earth would require modification to resolve most of global warming. This is highly dependent on the heat capacity and irradiance of the area selected. The versatile model was also used to provide UHIs global warming and cooling estimates.

## 1.0 Introduction

When we consider climate change solutions, in the race against time, it is advantageous to look at the practical aspects of implementing an albedo solution. Given the slow progress reported with greenhouse gas reduction, and the continual increase in the Earth’s average yearly temperature, it is important to revisit alternative albedo solutions. Unlike geoengineering solutions, GHG mitigation is highly difficult to result in reversing climate change, especially with reports on large deforestation occurring [1].

Implementation is a key focus on geoengineering an albedo surface solution. There have been a number of geoengineering solutions proposed [2-4] that are either atmospheric or surface-based. In this study, we focus on targeting surface regions and present practical engineering formulas and values.

The target areas that have the highest impacts are likely ones with:

- high solar irradiance
- large heat capacities
- low albedo
- ability to amplify nature’s albedo

To clarify the last target area, we infer that cooling down certain areas may cause natural compounding albedo changes to occur, such as increases in snowfall and ice formations. We can term these as Solar Amplified Areas (SAA) relative to Nominal Land Albedo (NLA) areas (25% albedo, see Sec. 7.2).

Although the task is highly challenging, it is easier to do geoengineering of surface reflectivity compared with building cities. Often, UHIs and impermeable surfaces are haphazardly constructed in terms of solar absorption considerations. While numerous authors [4-17] have found significant warming due to UHIs, the only motivated work in this area is a result of health concerns. Therefore, albedo cool roof solutions have not received adequate attention compared to GHG efforts. This is unfortunate and makes the business of solar solution and its financing less desirable. It is important that not just scientists understand the importance of the albedo solution. There is a lack of knowledge when it comes to the word albedo and its potential contribution. We cannot expect architects, road engineers, car designers, city planners, politicians and so forth, to do their job correctly in the green area, if these concepts are not widely understood. Therefore, a key strategy employed in this study is to demonstrate the advantages, feasibility and importance in cooling solar amplified areas made by man. We provide simple geoengineering equations that can aid the designer. We need to recognize that the whole is equal to the sum of the parts in GW; mankind’s resolve to greenhouse gases and albedo reductions, both need to be addressed for a realistic solution.

## 2. Outline of the Geoengineering the Albedo Solution

64 We present a brief outline to overview and clarify our modeling objectives and motivate interests.

65  
66 **Section 3:** In this section, we identify a practical re-radiation model to help obtain accurate important values in  
67 geoengineering a global warming albedo solution. In the absence of feedback, our GW model has the form:

$$68 \quad P_{Pre-Industrial} = P_{\alpha} + f_1 P_{\alpha} = \sigma T_S^4 \quad (1)$$

69  
70 Here  $P_{Pre-Industrial}$  is the total warming power (in  $W/m^2$ ),  $T_S$  is the Earth's average surface temperature,  
71  $P_{\alpha}=1361W/m^2/4 \times (1-\alpha)$  is the short wavelength absorption and  $f=\beta^4=0.618$  is a GHG re-radiation parameter, a  
72 redefined variable taken from the effective emissivity constant of the planetary system. The model is then extended  
73 so that it can be applied with climate feedback and verified using the Planck feedback parameter.

74  
75  
76 **Section 4:** Using the Model in Section 3, we apply it to temperature data from 1950 to 2019 and assess  $\Delta P_{Total}$ , the  
77 total forcing that has occurred. This is required in order to estimate the amount of reverse forcing corrective action  
78 needed.

79  
80 **Section 5:** In this section we first identify a key Planck-albedo parameter

$$81 \quad \gamma_{\% \Delta \alpha \Delta T} \approx 1W / m^2 / \Delta \%albedo / ^{\circ}K \quad (2)$$

82 The parameter converts a percent albedo  $\% \Delta \alpha$  change to  $\Delta P_T$ , the reverse forcing from the target area where the total  
83 reverse forcing  $\Delta P_{Rev\_S}$  is

$$84 \quad \Delta P_{Rev\_S} = -\gamma_{\% \Delta \alpha \Delta T} \% \Delta \alpha (1 + f_2) A_F = \Delta P_T (1 + f_2) A_F \quad (3)$$

85  
86 Here  $f_2$  is the 2019 re-radiation parameter, about 0.63 and  $A_F$  is an estimate of the anticipated GW feedback  
87 reduction.

88  
89  
90 **Section 6:** In this section an Albedo model is developed to use the  $\Delta P_T$  goal where

$$91 \quad \Delta P_T = \frac{A_T S_N}{A_E 4} 0.33 H_{T-N} [(\alpha'_T - \alpha_T)] \quad (4)$$

92 The factor,  $H_{T-N}$  is the hotspot irradiance sensible heat storage potential, a function of the heat capacity, mass,  
93 temperature storage, and solar irradiance by comparison to a nominal area. Here  $\alpha_T$  is the initial target albedo,  $\alpha'_T$   
94 is the modified target albedo, and 0.33 is the estimate fraction of time the target area is not covered by clouds. Then  
95 the final goal relative to fraction of Earth's area,  $A_E$ , needing modification is

- 96 •  $A_T / A_E$ , where  $A_T$  is the target area

97  
98  
99 **Section 7:** In this section, it all comes together by applying these models for different target areas including UHIs  
100 yielding their warming and cooling estimates.

101 Therefore, our task is to essentially find reasonable values for  $\Delta P_{Total}$ ,  $f_2$ ,  $\Delta P_{Rev\_S}$ ,  $H_{T-N}$ ,  $\gamma_{\% \Delta \alpha \Delta T}$ ,  $A_F$ ,  $\Delta P_T$ ,  $\% \Delta \alpha$ , in  
102 order to estimate a geoengineering GW solution by modifying the select fractional target area  $A_T/A_E$  of the Earth.

### 103 3.0 The Re-radiation Global Warming Model

104 In geoengineering, we are working with absorption and re-radiation, we define

$$105 \quad P_{Total} = \sigma T_S^4 = \sigma \left( \frac{T_e}{\beta} \right)^4 \quad \text{and} \quad P_{\alpha} = \sigma T_{\alpha}^4 = \sigma (\beta T_S)^4 \quad (5)$$

106 The definitions of  $T_{\alpha}=T_e$ ,  $T_S$  and  $\beta$  are the emission temperature, surface temperature and  $\beta=0.887$ , respectively.  
107 Consider a time when there is **no feedback issues** causing warming trends. Then by conservation of energy, the  
108 equivalent power re-radiated from GHGs in this model is dependent on  $P_{\alpha}$  with

$$109 \quad P_{GHG} = P_{Total} - P_{\alpha} = \sigma T_S^4 - \sigma T_{\alpha}^4 \quad (6)$$

110 To be consistent with  $T_{\alpha}=T_e$ , since typically  $T_{\alpha} \approx 255^{\circ}K$  and  $T_S \approx 288^{\circ}K$ , then in keeping with a common definition of  
111 the global beta (the proportionality between surface temperature and emission temperature) for the moment  
112  $\beta=T_{\alpha}/T_S=T_e/T_S$ .

121  
122 This allows us to write the dependence  
123

$$124 \quad P_{GHG} = \sigma T_S^4 - \sigma T_\alpha^4 = \frac{\sigma T_\alpha^4}{\beta^4} - \sigma T_\alpha^4 = \sigma T_\alpha^4 \left( \frac{1}{\beta^4} - 1 \right) = \sigma T_\alpha^4 \left( \frac{1}{f} - 1 \right) \quad (7)$$

125  
126 Note that when  $\beta^4=1$ , there are no GHG contributions. We note that  $f$ , the re-radiation parameter equals  $\beta^4$  in the  
127 absence of feedback.

128  
129 We can also define the blackbody re-radiated by GHGs given similarly by some fraction  $f_1$  such that  
130

$$131 \quad P_{GHG} = f_1 P_\alpha = f_1 \sigma T_\alpha^4 \quad (8)$$

132  
133 It is important in geoengineering to view the re-radiation as part of the albedo effect. This is a key difference in how  
134 we view the total effect from short wavelength absorption by the inclusion of the re-radiation effect. Consider  $f=f_1$ ,  
135 in this case according to Equations 7 and 8, it requires  
136

$$137 \quad P_{GHG} = \sigma T_\alpha^4 \left( \frac{1}{f} - 1 \right) = f_1 \sigma T_\alpha^4 = f \sigma T_\alpha^4 \quad (9)$$

138  
139 This dependence leads us to the solution of the quadratic expression  
140

$$141 \quad f^2 + f - 1 = 0 \text{ yielding } f = 0.618034 = \beta^4, \beta = (0.618034)^{1/4} = 0.88664 \quad (10)$$

142  
143 This is very close to the common value estimated for  $\beta$  and this has been obtained through energy balance in the  
144 planetary system providing a self-determining assessment. In Appendix A , we double check this model in another  
145 way by balancing energy in and out of our global system. Then in Section 4.2, we apply the model to demonstrate its  
146 capability and consistency with the Planck parameter. We note that the assumption  $f=f_1$  only works if planetary  
147 energy is in balance (also see Appendix A) without feedbacks.

#### 149 4.0 Re-radiation Model Applied to Two Different Time Periods

150  
151 Global warming can be exemplified by looking at two different time periods. The model applied for 1950 needs to  
152 be consistent with Eq. 6 and 8. Here we will  
153

- 154 • assume no feedback issues causing a warming trend in 1950 so that from our model

$$155 \quad P_{Total,1950} = P_\alpha + P_{GHG} = P_\alpha + f_1 P_\alpha = P_\alpha (1 + f_1) = 1.618 P_\alpha \quad (11)$$

156  
157 where  $P_\alpha = S_o \{0.25x(1 - Albedo)\}$  and  $S_o=1361W/m^2$ . Although 1950 is not truly pre-industrial (see Eq. 1), we  
158 proceed under the assumption of no changes in GHG and feedback issues at this time to establish our baseline, since  
159 geoengineering a solution to earlier dates would pose even higher challenges. Under this assumption,  $1+f=1.618$   
160 becomes the 1950 albedo-GHG reference value. Since its value is related to the re-radiation parameter, it is  
161 subjected to changes due to variations in our aging climate system. As a reference value, it is constrained by the  
162 energy balance in Eq. 9 and as discussed in Section 4.2.  
163

164  
165 In 2019 due to global warming trends, this model is more complex and harder to separate out terms. However, we  
166 are still able to obtain reasonable accuracy since we are fitting the model to the Earth's average temperature data  
167 with the goal in mind of finding reasonable accurate estimate of total forcing  $\Delta P_{Total} = P_{Total,1950} - P_{Total,2019}$ . Therefore,  
168 we proceed similarly and results and verification will also justify its use, then  
169

$$170 \quad P_{Total,2019} = P_{\alpha'} + P_{GHG'+Feedback} \approx P_{\alpha'} + f_2 P_{\alpha'} \quad (12)$$

171  
172 Here,  $P_{GHG'+Feedback}$  includes the 1950 GHGs and 2019 increase with feedbacks such as water-vapor concentration,  
173 lapse rate effect and other changes including increase in snow-ice albedo variations that are hard to separate out.  
174 That is, feedbacks are related to GHG increases and albedo change.  $P_{\alpha'}$  represents the 2019 point in time with its  
175 albedo due to changes in UHI absorption, cloud absorption, ice and snow melting, and so forth that can be discerned.  
176 The model does not demand rigid accountability in its application (see Sec.4.2) but reasonable estimates are helpful.  
177 We note that unlike  $f_1$ ,  $f_2$  is not a strict measure of the emissivity.

178  
179 In 1950  $f_1$  defines the GHG re-radiation function (with no feedbacks) and is consistent with the estimates for beta. In  
180 2019, it is more complex and according to Eq. 12, must include feedbacks. The value  $f_2$  while close to the beta value  
181 in Eq. 10, is no longer identical as  $f_1$  (see Equation 13). The value  $f_2$  can also be assessed relative to  $f_1$  as described  
182 in the next section. However, in general, between the two time periods, we will find  $P_{GHG} \approx P_{GHG'+Feedback}$  (see results  
183 in Section 4.2).

#### 184 185 **4.1 Warming Imbalance in 2019**

186  
187 The re-radiation parameters  $f_1$  and  $f_2$ , are connected and from Eq. 10, 11 and 12 we have  
188

$$189 \quad f_2 = f_1 + \left( \frac{P_{2019}}{P_{\alpha'}} - \frac{P_{1950}}{P_{\alpha}} \right) = f_1 + \left( \frac{P_{GHG'+F}}{P_{\alpha'}} - \frac{P_{GHG}}{P_{\alpha}} \right) = f_1 + \Delta f = \beta_1^4 + \Delta f \approx \beta_2^4 + \Delta f \quad (13)$$

190 In this way  $f_2$  is a function of  $f_1=0.618$  and the differences in the global warming residuals that is identified in Eq. 12  
191 as  $\Delta f$ . The RHS of Eq. 12 (indicating that  $\beta_1 \approx \beta_2$ ) will become apparent in application (Eq. 16 and 17) and  
192 verification.

#### 193 194 **4.2 Results Applied to 1950 and 2019**

195  
196 Since the re-radiation parameter is fixed for  $f_1=0.618$ , to obtain the average  $T_{1950}=13.89^\circ\text{C}$  ( $287.038^\circ\text{K}$ ), the only  
197 adjustable parameter left in our model is the global albedo. This requires an albedo value of 0.3008 (see Table 1) to  
198 obtain the correct value  $T_{1950}$ . This albedo number is reasonable and similar to values cited in the literature [18].  
199

200 In 2019, the average temperature of the Earth is  $T_{2019}=14.84^\circ\text{C}$  ( $287.99^\circ\text{K}$ ). Here we are not sure of the albedo value  
201 since it likely changed due to UHI increase, snow and ice melting and cloud coverage changes. The IPCC value in  
202 AR5 [19] is 0.294118 (100/340). However, this would represent a 3% change since 1950 which may be an  
203 overestimation. In this assessment, we will assume a low middle value of 1.2% change. Another reason for this  
204 choice is in a resulting analysis in Appendix A.2. Then, the  $f_2$  parameter is adjusted to 0.6311 to obtain  $T_{2019}$ . Table  
205 1 summarizes model results for the specified albedos and observed Earth's surface temperatures. The results yield  
206  $P_{Total\ 1950}=384.935\ \text{W/m}^2$  and  $P_{Total\ 2019}=390.055\ \text{W/m}^2$ .  
207  
208

**Table 1** Model results

| Year               | $T(^{\circ}\text{K})$ | $T_{\alpha}(^{\circ}\text{K})$ | $f_1, f_2$    | $\alpha, \alpha'$ | $P_{\alpha}, P_{\alpha'}$<br>( $\text{W/m}^2$ ) | $P_{GHG'+feedback}$<br>$P_{GHG}$ ( $\text{W/m}^2$ ) | $P_{Total}$<br>( $\text{W/m}^2$ ) |
|--------------------|-----------------------|--------------------------------|---------------|-------------------|---|---|-----------------------------------|
| 2019               | 287.991               | 254.83                         | 0.63114       | 29.719            | 239.131   | 150.925   | 390.056                           |
| 1950               | 287.041               | 254.51                         | 0.6180        | 30.08             | 237.903   | 147.032   | 384.935                           |
| $\Delta 2019-1950$ | <b>0.95</b>           | 0.328                          | <b>1.311%</b> | 0.361             | <b>1.228</b>                                    | 3.893   | <b>5.12</b>                       |
|                    |                       |                                |               | <b>(1.2%)</b>     |   |   |                                   |

209  
210 From Table 1 we now have identified the reverse forcing at the surface needed since  
211

$$212 \quad \Delta P_{Total} = P_{2019} - P_{1950} = 5.121\ \text{W} / \text{m}^2 \quad (14)$$

213  
214 and

$$215 \quad \Delta T_{Total} = T_{2019} - T_{1950} = 0.95^\circ\text{C} \quad (15)$$

216  
217 as modeled.

#### 218 219 **4.3 Showing Model Consistency with the Planck Parameter**

220  
221 To show model consistency, the forcing change,  $5.121\ \text{W/m}^2$ , resulting in a  $0.95^\circ\text{K}$  rise, should agree with what is  
222 expected when using the Planck feedback parameter.  
223

224 In order to show model consistency, we will need some exact values for beta using the temperatures in Table 1,  
225 these are from the two different time periods (see Eq. A-3)  
226

$$227 \quad \beta_{1950} = \frac{T_{\alpha}}{T_s} = \frac{T_e}{T_s} = \frac{254.51}{287.04} = 0.88667 \text{ and } \beta_{1950}^4 = 0.61809 \quad (16)$$

228  
229 as this value is consistent with Eq. 10, and  
230

$$\beta_{2019} = \frac{T_e}{T_s} = \frac{T_e}{T_s} = \frac{254.83}{287.99} = 0.88485 \text{ and } \beta_{2019}^4 = 0.61304 \quad (17)$$

232  
233 Although these two are very close, we use both values due to the need for high accuracy; model self-consistency is  
234 required.

235  
236 From the definition of the Planck parameter and results in Table 1, we can estimated [20]  
237

$$\lambda_o = -4 \frac{\Delta R_{LWR}}{T_s} = -4 \left( \frac{237.9W/m^2}{287.04^\circ K} \right)_{1950} = -3.315W/m^2/^\circ K \quad (18)$$

239 and

$$\lambda_o = -4 \frac{\Delta R_{LWR}}{T_s} = -4 \left( \frac{239.13W/m^2}{287.99^\circ K} \right)_{2019} = -3.321W/m^2/^\circ K \quad (19)$$

241  
242 We note these are very close in value showing minor error and consistency with Planck parameter value, often taken  
243 as  $3.3W/m^2/^\circ K$ . While there are only small differences between each beta and these two Planck parameters, final  
244 warming predictions using a Planck parameter method, requires values found from the model. This self-consistency  
245 helps in providing accuracy for estimating  $\Delta T$  by reducing compounding error within the model. We then use the  
246 generalized form for the long wavelength estimate in Equation A-2, yielding the approximate warming change in  
247 terms of the total power and the Planck parameter method as [20]  
248

$$\Delta T = T_{1950} - T_{2019} = -4 \left\{ \left( \frac{\beta^4 P_{Total}}{\lambda_o} \right)_{1950} - \left( \frac{\beta^4 P_{Total}}{\lambda_o} \right)_{2019} \right\} \quad (20)$$

250  
251 Using Table 1, the temperature warming results is  
252

$$\Delta T = -4 \left( \frac{0.6181 \times 384.935W/m^2/^\circ K}{3.315W/m^2/^\circ K} - \frac{0.61304 \times 390.056W/m^2/^\circ K}{3.3215W/m^2/^\circ K} \right) = 0.92^\circ K \quad (21)$$

254  
255 This equation illustrates consistency of the re-radiation model with the Planck parameter showing reasonable  
256 accuracy helping to verify the model from a different perspective. The model allows for a number of helpful  
257 comparisons that are described in Appendix A.2.  
258

## 259 5.0 Geoengineering Reverse Forcing Solution

260  
261 The albedo changes and  $\Delta P_\alpha$  in Table 1, are:  $\% \Delta \alpha = 1.2\%$  and  $1.228W/m^2$ , respectively. We note that we can define  
262 a unique Planck-albedo parameter  $\gamma_{\% \Delta \alpha} = \Delta P_\alpha / \% \Delta \text{albedo}$ . To illustrate from Table 1

$$\gamma_{\% \Delta \alpha} = 1.023 W/m^2/\Delta \% \text{albedo} \quad (22)$$

266 This parameter can also be expressed per degree (noting the  $0.95^\circ K$  change in Table 1)  
267

$$\gamma_{\% \Delta \alpha \Delta T} \approx 1W/m^2/\Delta \% \text{albedo}/^\circ K \quad (23)$$

269  
270 The helpful parameter [5] is featured here as a modeling tool. We term it the Planck-albedo parameter, since it  
271 relates to blackbody ( $P_\alpha$ ) absorption. A simple numeric example is given in Sec. 5.1 to illustrate how it provides  
272 helpful estimates along with the albedo-GHG factor. This interesting parameter simplifies and is not really  
273 dependent on two different time periods estimates of the global alpha changes since  
274

$$\gamma_{\% \Delta \alpha} = \frac{(\Delta E_o)_\alpha}{\frac{\alpha_1 - \alpha_2}{\alpha_1} 100} = \frac{E_o (\alpha_1 - \alpha_2)}{\frac{\alpha_1 - \alpha_2}{\alpha_1} 100} = E_o \alpha_1 / 100 \approx 1W/m^2/\% \Delta \text{albedo} \quad (24)$$

276  
277 where  $E_o = 340 W/m^2$  and when  $\alpha_1$  is 0.294118, the value  $1.000W/m^2/\Delta \% \text{albedo}$  is obtained. We note the value  
278  $29.4118\%$  ( $100/340$ ) is given in AR5 [19].  
279

280 The albedo-GHG and the Planck-Albedo parameter may now be combined in order to provide a simple solar  
281 geoengineering solution estimate for reverse forcing

$$\Delta P_{\text{Rev}_S} = -\gamma_{\% \Delta \alpha \Delta T} \% \Delta \alpha (1 + f_2) A_F = \Delta P_T (1 + f_2) A_F \quad (25)$$

These variables have been defined in the outline (Section 2.0). This equation provides a fairly simple and practical way to estimate  $\Delta P_{\text{Rev}_S}$ . In solar geoengineering, anticipating an allowance for the climate system to equilibrate [21] is not considered here. Furthermore, one might expect that a positive compared to negative albedo change may not have a strong hysteresis effect (as long as a tipping point has not occurred). Note that the  $1+f$  factor accounts for one process of initial absorption change  $\Delta P_T$  followed by subsequent partial re-radiation from GHGs. This value helps to clarify our goal.

The effective results

$$\text{Effect} = \frac{\Delta P_{\text{Rev}_S}}{\Delta P_{\text{Total}}} \quad (26)$$

and  $\Delta P_{\text{Rev}_{OLWR}} = \beta^4 \Delta P_{\text{Rev}_S}$  the temperature reduction can be estimated from [20]

$$\Delta T_{\text{Rev}} = -\frac{\beta^4 \Delta P_{\text{Rev}_S}}{\lambda_o} \quad (27)$$

In theory,  $\Delta T_{\text{Rev}}$  is only an estimate since this equation is valid when no feedback issues result. The reason it is a reasonable estimate is that  $\beta^4 \Delta P_{\text{Rev}_S}$  is a good estimate OLWR (also see Eq. A-2).

### 5.1 Example of a Reverse Forcing Goal

In this section, we consider a goal of 1.5% geoengineering albedo change. Using Equation 25, with a decrease in water-vapor feedback anticipated, we might use a value of  $A \approx 2$  [21], then

$$\Delta P_{\text{Rev}_S} = -1 \text{ W/m}^2 / \% \times 1.5\% \times (1 + f_2) \times 2 = -1.5 \text{ W/m}^2 \times (1 + f_2) \times 2 = -4.9 \text{ Watt/m}^2 \quad (28)$$

This estimate can be compared with the re-radiation model results in Table 1 showing a forcing of  $5.21 \text{ W/m}^2$  to obtain the relative effect of 94% from Eq. 26 for this particular geoengineering solution. From Equation 28 an estimate of the temperature cooling can be obtained with Eq. 27 where  $\beta^4 \Delta P_{\text{Rev}_S} = \Delta P_{\text{Rev}_{OLWR}} = -3.0 \text{ W/m}^2$  is

$$\Delta T_{\text{Rev}} = \frac{3.0 \text{ W/m}^2}{\lambda_o} = -0.91^\circ \text{K} \quad (29)$$

This would indicate a significant resolution to the current warming trend. As one might suspect, a 1.5% albedo change requires a lot of modified area. Feasibility is discussed in the rest of this paper. We note a number of solar geoengineering solutions have been proposed [2-4].

## 6.0 Converting the Reverse Forcing Goal to Target Area

We can write the short wavelength solar absorption as

$$P = \frac{Q}{A} = \frac{S_N}{4} \sum_i \frac{A'_i}{A} (1 - \alpha_i) + \frac{S_N}{4} H_{T-N} \frac{A'_T}{A} (1 - \alpha_T) + \frac{S_N}{4} \frac{A_C}{A} (1 - \alpha_C) \quad (30)$$

Here  $A_i$  is the  $i^{\text{th}}$  effective area having an albedo  $\alpha_i$ ,  $S_N = 1361 \text{ W/m}^2$  and  $A$  is the surface area of the Earth and  $A_C$  is effective cloud coverage. We consider a change to a hotspot target effective area  $A_T$  with albedo  $\alpha_T$ . In addition, because we select a particularly problematic solar absorbing target compared to a nominal area (N), it has hotspot irradiance sensible heat storage potential  $H_{T-N}$ , a function of the heat capacity, mass, temperature storage, and solar irradiance. Essentially this has the effect of amplifying the target area.  $H_{T-N}$  is described and enumerated in Appendix B.

We note that the Earth Albedo change will only be a function of the target area variation, so from Eq. 30

$$(dP_T)_\alpha = \frac{S_N}{4} H_{T-N} \frac{A'_T}{A} (-d\alpha_T) \quad (31)$$

where the subscript  $\alpha$  indicates all other Earth albedo components are held constant.

334 The overall equation prior to changing the albedo is subject to the constraints

335

$$336 \quad P = 240W / m^2 \text{ and } A = \sum_i A'_i + A_T' + A_C = A_E' + A_C \text{ but } A_E' = (1 - \%A_C)x A_E = 0.33A_E \quad (32)$$

337

338 This indicates that because of the cloud coverage term  $A_C$ , about 67% of the actual Earth's area  $A_E$  [23] is covered  
339 from direct sunlight. This is likely conservative as clouds do let some sunlight through. However, that means that  
340 roughly 33% of the time areas receive sun during daylight hours.

341

342 We now alter the target albedo  $\alpha_T$  to  $\alpha_T'$  of a SAA so that

343

$$344 \quad P' = \frac{Q'}{A} = \frac{S_N}{4} \sum_i \frac{0.33A_i}{A} (1 - \alpha_i) + \frac{S_N}{4} \frac{0.33A_T}{A} H_{T-N} (1 - \alpha_T') + \frac{S_N}{4} \frac{A_C}{A} (1 - \alpha_C) \quad (33)$$

345

346 Note the 0.33 factor is now added due to the percent of time the albedo change is effective. Using the example goal  
347 of the target area  $\Delta P_T = 1.5W/m^2$  in Eq. 28, the change in heat absorbed is a function of the target area as indicated by  
348 Eq. 31, where

$$349 \quad \Delta P_T = P - P' = \frac{S_N}{4} \frac{0.33A_T H_{T-N}}{A} [(\alpha_T' - \alpha_T)] = 1.5W / m^2 \quad (34)$$

350

351 However, the same results can be obtained by changing the albedo of a nominal area; so in this case  $H_{T-N} = 1$ . The  
352 equivalent change for the NLA is

353

$$354 \quad \Delta P_{T-N} = \frac{S_N}{4} \frac{0.33A_N}{A} \{(\alpha_N' - \alpha_N)\} = 1.5W / m^2 \quad (35)$$

355

## 356 7.0 Geoengineering Application

357

358 Comparing the target to the nominal areas, we have

359

$$360 \quad \frac{\Delta P_T}{\Delta P_{T-N}} \approx \frac{A_T H_{T-N} [(\alpha_T' - \alpha_T)]}{A_N [(\alpha_N' - \alpha_N)]} = 1 \quad (36)$$

361

362 As an example, assume  $H_{T-N} \approx 9$  (see Appendix B),  $\alpha_N = 0.25$  (see Sec. 7.2),  $\alpha_T = 0.12$  [24], and for  $\alpha_N' = \alpha_T' = 0.9$ , we  
363 obtain

$$364 \quad \frac{A_N}{A_T} = \frac{H_{T-N} [(\alpha_T' - \alpha_T)]}{[(\alpha_N' - \alpha_N)]} = \frac{9[(0.9 - 0.12)]}{[(0.9 - 0.25)]} = 10.8 \quad (37)$$

365

366 This indicates that the nominal area would have to be about 11 times larger than the target area for equivalent  
367 results.

368

369 In assessing our goal, we have from Eq. 28

370

$$371 \quad \Delta P_T = \frac{S_N}{4} \frac{0.33A_T H_{T-N}}{A_E} [(\alpha_T' - \alpha_T)] = 1.5W / m^2 \quad (38)$$

372

373 For  $H_{T-N} = 1$ ,  $\alpha_T' = 0.9$ , and  $\alpha_T = 0.12$  then

374

$$375 \quad \Delta P_T = 340 \frac{A_T}{A} [0.78] x 0.33 = 1.5W / m^2 \quad (39)$$

376 and

$$377 \quad \frac{A_T}{A} = 0.01714 = 1.714\% \text{ of Earth} \quad (40)$$

378

379 For  $H_{T-N}=10$ ,  $\alpha_T'=0.9$ , and  $\alpha_T=0.12$  then

380

$$381 \quad \frac{A_T}{A} = 0.1714\% \text{ of Earth} \quad (41)$$

382 Recall that the goal for a  $1.5W/m^2$  corresponded to a 1.5% albedo change (see Sec. 5.1). We can check results of  
 383  $A_T/A=1.714\%$  when  $H_{T-N}=1$ , yields a 1.5% albedo change using a related expression to Eq. 38. This is given by

384

$$385 \quad \Delta\alpha\% = 0.33 \frac{A_T}{A} \frac{[(\alpha_T' - \alpha_T)]}{\alpha} = 0.33(1.714\%) \frac{[(0.9 - 0.12)]}{0.294118} = 1.5\% \quad (42)$$

386

387 where the global albedo is taken as  $\alpha=0.294118$  which is indicated in AR5's energy budget figure [19].

388

### 389 7.1 Cooling Estimates Compared to Urban Heat Island Area

390

391 Since UHI are likely good target areas, we can compare these results to the total global urbanized area. Such  
 392 estimates of urbanization unfortunately vary widely partly due to the confusing definition of what is urban.  
 393 However, two studies are of interest. A Schneider study [25] on 2000 data estimated that 0.148% of the Earth was  
 394 covered by UHI and the associated surrounding urban areas. Due to city growth, this extrapolates to 0.188% [5] in  
 395 2019. Similarly, a study from GRUMP [26] showing global urbanization value in 2000 of 0.783% extrapolates to  
 396 0.953% [5] of the Earth's area in 2019. These extrapolations are based on an average yearly urbanization growth  
 397 rate between 1.3% to 1.6% [5]. Lastly, note that UHIs have their own hotspot amplification factors [5] that vary  
 398 between 3.1 and 8.4 (see Appendix C) which are listed in Table 2 and can be applied for  $H_{T-N}$ . Therefore, compared  
 399 to these 2019 estimates for urban heat island and surrounding areas, the required area changes for different  $H_{T-N}$   
 400 values (discussed in Appendix C) are summarized in Table 2.

401

402

**Table 2** Cooling required areas relative to UHI areas

| $H_{T-N}$ | $A_T/A$<br>(% of<br>Earth)<br>$\alpha_T' = 0.9$ | Schneider Factor<br>( $A_T/A$ )/0.188%<br>(Conservative)<br>$\alpha_T' = 0.9$ ( $\alpha_T' = 0.5$ ) | GRUMP Factor<br>( $A_T/A$ )/0.953<br>$\alpha_T' = 0.9$ ( $\alpha_T' = 0.5$ ) |
|-----------|---|---|--|
| 1         | 1.714   | 9.1 (18.7)  | 1.8 (3.7)  |
| 3.1       | 0.55  | 2.93 (6)  | 0.58 (1.2)   |
| 8.4       | 0.2   | 1.06 (2.2)  | 0.21 (0.43)  |
| 9         | 0.19  | 1 (2.1)   | 0.2 (0.41)   |

403

\* $A_T/A$  represent 94% of the solution (see Sec. 5.1)

404

405 Note that an IPCC (Satterthwaite et. al. [27]) AR5 report references the Schneider et al. [25] results in urban  
 406 coverage of 0.148% of the Earth.

407

408 Table 2 results are highly dependent on target albedo change and  $H_{T-N}$  which is overviewed in Appendix B. It is  
 409 important to develop better estimates for both  $H_{T-N}$  and urbanization sizes than estimated here. We note that the 0.12  
 410 albedo value applies to UHI [24], which is acceptable upper value when looking for hotspot targets. The albedo and  
 411 two  $H_{T-N}$  values cited here have been studied in Feinberg [5]. The assessments for  $H_{T-N}$  applicable to UHIs are also  
 412 provided to aid the reader in Appendix C. Results in Table 2 illustrate feasibility and the probable geoengineering  
 413 challenges. A worldwide effort would provide motivation from a number of key benefits; resolving much of global  
 414 warming, providing assurance against a tipping point, and local health benefits by cooling off cities. UHIs pose a  
 415 number of challenges in trying to cool off their areas. The Schneider results in row 2, indicate that the potential area  
 416 needed may be 3-6 times their current size. Therefore, if this was proven to be the most accurate estimate,  
 417 supplementary target areas would be required to reach the 94% objective. Furthermore it is unrealistic to realize an  
 418 overall UHI albedo goal of 0.9 due to their complex nature so we have also provided goals at 0.5 as well in the table.

419

420 Generally, UHIs meet a lot of the requirements for good targets having high heat capacity with large hotspot areas  
 421 and massive sensible heat storage. One helpful aspect to note is that cool roof implementation also allows for more  
 422 stable albedo maintenance over time compared to other areas like mountain regions. However, the complex nature  
 423 of cities also makes it highly challenging.



424

425 **7.2 Warming Estimates Due to Urban Heat Island Area**

426

427 We can use this same model to estimate the global warming contributions due to UHIs. In this case, instead of  
 428  $\alpha_T'=0.9$ , we evaluate by restoring the UHIs to their original estimated albedo value of  $\alpha_T'=0.25$ . This albedo value is  
 429 based on a study by He et. al. [28] which found the land albedo varied from 0.1 to .4 having an average of 0.25.  
 430 Then using the  $H_{T-N}$  values in Section 7.1, we estimate the percent of the Earth needed to obtain a 94% solution and  
 431 compare results to the known UHI coverage areas.

432

433 For  $H_{T-N}=3.1$ ,  $\alpha_T'=0.25$ , and  $\alpha_T=0.12$  then

434

$$435 \Delta P_T = \frac{1361W/m^2}{4} \frac{0.33A_T 3.1}{A_E} [(0.25 - 0.12)] = 1.5W/m^2 \quad (43)$$

436 and

$$437 \frac{A_T}{A} = 3.3\% \quad (44)$$

438

439 of the Earth. Similarly for  $H_{T-N}=8.4$ ,  $\alpha_T'=0.25$ , and  $\alpha_T=0.12$  then

440

$$441 \frac{A_T}{A} = 1.2 \% \text{ of Earth} \quad (45)$$

442

443 Table 3 summarized the warming trend results

444

445

**Table 3** UHI Warming estimates

| $H_{T-N}$ | $A_T/A$<br>(% of<br>Earth) | Schneider Factor<br>( $A_T/A$ )/0.188%<br>(Conservative) | GRUMP Factor<br>( $A_T/A$ )/ 0.953 | GW%<br>1/Schneider<br>Factor<br>/ 0.94* | GW%<br>1/GRUMP<br>Factor<br>/ 0.94* |
|-----------|----------------------------|--|------------------------------------|---|-------------------------------------|
| 3.1       | 3.3                        | 17.6   | 3.5                                | 6.1                                     | 31                                  |
| 8.4       | 1.2                        | 6.4  | 1.26                               | 16.9                                    | 85.4                                |

446 \* $A_T/A$  GW represent 94% of the solution (see Sec. 5.1), and are adjusted to 100% in Column 5 & 6

447

448 Results in Column 5 and 6 are reasonably comparable to Feinberg 2020 [5]. The model shows that between 6.1%  
 449 and 85% of global warming could be due to UHIs and their coverage. We note these large variations are due to the  
 450 difficulty in estimating  $H_{T-N}$  and knowledge of UHI area coverages, as shown in the differences found between  
 451 Schneider and the GRUMP studies. However, the model provides a reasonable way to make estimates which can be  
 452 further refined once better values are known.

453

454 Furthermore, we note the cooling potential in Table 2 is about a factor of 3 to 6 times compared to the warming  
 455 shown in Table 3. For example in Table 2 and 3, the area full warming to cooling ratio 17.6/2.93 yields an effective  
 456 potential factor of 6 for  $\alpha_T'=0.9$ , and a factor of 2.9 (17.6/6) for  $\alpha_T'=0.5$ . As stated above, obtaining the full cooling  
 457 potential ( $\alpha_T'=0.9$ ) for UHIs and their impermeable surfaces is likely unobtainable due to the complex nature of  
 458 cities therefore the value  $\alpha_T'=0.5$  is a better guide.

459

460 **7.3 Some Hotspot Target Areas**

461

462 There are many hotspots that provide likely target areas. Deserts would be highly difficult to maintain any albedo  
 463 change. However, mountains, UHI cool roofs in cities, and impermeable surface such as roads might be logical  
 464 target areas. Some interesting known hotspots include

465

- 466 • Flaming Mountains, China
- 467 • Bangkok, Thailand (planet's hottest city)
- 468 • Death Valley California
- 469 • Titat Zvi, Israel

- 470 • Badlands of Australia
- 471 • Urban Heat Islands & all Impermeable surfaces
- 472 • Oceans [2]

473  
 474 We note that mountain areas in cool regions should not be excluded; natural compounding albedo effects may occur  
 475 from increases in snow-fall and ice formations. Albedo changes could be performed in summer months and then in  
 476 winter months compounding effects assessed.

477  
 478 As a summary, Equations 25 and 35 can be combined to provide a resulting solar geoengineering equation for  
 479 reverse forcing obtained in this study where

$$481 \quad \Delta P_{\text{Rev}_S} = -\gamma_{\% \Delta \alpha \Delta T} \% \Delta \alpha (1+f) A_R = \left\{ \frac{S_N}{4} 0.33 H_{T-N} \frac{A_T}{A} [(\alpha'_T - \alpha_T)] \right\} (1+f) A_R \quad (46)$$

482  
 483 with suggested values  $H_{T-N}=6$ ,  $\alpha'_T=0.9$ ,  $\alpha_T=.12$ ,  $\Delta P_{\text{Rev}_S}=4.8\text{W/m}^2$ , and  $f=0.63$ .

## 484 485 8. Conclusions

486  
 487 The albedo solution is vital in mitigating global warming. Today, technology has numerous advances that include  
 488 improvements in materials, drone capability, artificial intelligence, which could be helpful in geoengineering  
 489 surfaces. Mankind has addressed many technological challenges successfully. It is not illogical to consider a global  
 490 albedo solution while time permits prior to a potential tipping point.

491  
 492 In this paper we have provided a number of important estimates that include:

- 493
- 494 • A target albedo goal of  $-4.8\text{W/m}^2$  ( $\Delta P_{\text{Rev}_LWR}=-2.97\text{W/m}^2$ )
- 495 • The target area required to resolve 94% of global warming is about 0.2% to 0.5% (Table 2) of the Earth, if  
 496 proper hotspots are cooled with highly reflective surfaces. This is likely on the order of UHIs coverage  
 497 today
- 498 • The cooling potential of UHIs is about a factor of 3 time higher than their warming contribution if highly  
 499 reflective surfaces can be realized
- 500 • Likely target areas may include problematic hotspots such as UHIs, mountains regions and possibly ocean  
 501 areas [2]
- 502 • Selecting proper hotspots can reduce the required target area by an estimated factor of 11
- 503 • Changing the albedo has 1.6 benefit factor due to GHG re-radiation
- 504 • UHIs likely contribute significantly to global warming
- 505 • Solutions are highly dependent on  $H_{T-N}$ .

506  
 507 Finally we suggest:

- 508
- 509 • Tasking agencies worldwide, such as NASA, to work full time on solar geoengineering, which at this late  
 510 time should be our highest priority,
- 511 • Worldwide albedo guidelines for both UHIs and impermeable surfaces similar to on-going  $\text{CO}_2$  efforts
- 512 • Worldwide guidelines for future albedo design considerations of cities,
- 513 • Changing impermeable surfaces of roads, sidewalks, driveways, parking lots, industrial areas such as  
 514 airports, distribution centers, and roof tops to reflective surfaces. We note that their cooling potential can be  
 515 much larger compared to their warming contribution, and a full review should be performed. Furthermore,  
 516 such surfaces create hydro-hotspots [29] which may contribute to higher values of  $H_{T-N}$ . A hydro-hotspot is  
 517 a hot surface that creates moisture in the presence of precipitations. Such surfaces create excess moisture in  
 518 the atmosphere promoting a local greenhouse effect.
- 519 • Manufacturing cars to be more reflective including reducing their internal solar heating. Although,  
 520 worldwide solar cool vehicles (e.g., silver or white) will likely not contribute significantly to global  
 521 warming mitigation, recommending them would. It will help raise awareness, similar to electric  
 522 automobiles that help improve  $\text{CO}_2$  emissions and could increase interest in similar projects thereby  
 523 promoting other related changes like cool roofs.

## 524 Appendix A: Re-radiation Model's Energy Balance

526  
527 Although  $f_1$  has been uniquely defined in Eq. 10, this should also result from balancing the energy in and out of the  
528 global system.

### 529 530 A.1 Balancing $P_{out}$ and $P_{in}$ in 1950

531  
532 To balance the energy in 1950, we start with Eq. 11. In equilibrium the radiation that leaves must balance  $P_{\alpha}$ , from  
533 the energy absorbed, so that

$$534 \quad \begin{aligned} \text{Energy}_{out} &= (1-f_1)P_{\alpha} + (1-f_1)P_{Total} = (1-f_1)P_{\alpha} + (1-f_1)\{P_{\alpha} + f_1P_{\alpha}\} \\ &= 2P_{\alpha} - f_1P_{\alpha} - f_1^2P_{\alpha} = \text{Energy}_{in} = P_{\alpha} \end{aligned} \quad (\text{A-1})$$

536  
537 This is consistent, so that in 1950 Eq. A-1 requires the same quadratic solution as Eq. 10. It is also apparent that  
538

$$539 \quad P_{\alpha} = f_1 P_{Total\_1950} = \beta_1^4 P_{Total\_1950} \quad (\text{A-2})$$

540 since

$$541 \quad P_{\alpha} = f_1(P_{\alpha} + f_1P_{\alpha}) \quad \text{or} \quad 1 = f_1(1 + f_1) \quad (\text{A-3})$$

544  
545 The RHS of Eq. A-3 is Eq. 10. This illustrates  $f_1$  from another perspective as the fractional amount of total radiation  
546 in equilibrium. As a final check, the application in Section 4.2, Table 1, illustrate that  $f_1$  provides reasonable results.

### 547 548 A.2 Comparisons Using the Albedo-GHG Factor

549  
550 We can look at an important ratio, the power created by the albedo effect compared to GHGs in 1950. The initial  
551 radiation is  $P_{\alpha}$ , and then according to Eq. 11 and Table 1, the energy is increased by  $P_{GHG}$  due to re-radiation  $fP_{\alpha}$  that  
552 yields the ratio

$$553 \quad \left\{ \frac{P_{\alpha} + P_{GHG}}{P_{GHG}} = \frac{P_{\alpha} + f_1P_{\alpha}}{f_1P_{\alpha}} = \frac{1+f_1}{f_1} = \frac{1.62}{0.62} = 2.62 \right\}_{1950} \quad \text{also note that} \quad \left\{ \frac{1+f_2}{f_2} = 2.58 \right\}_{2019} \quad (\text{A-4})$$

555  
556 We note the ratio is reduced in 2019 due to the addition  $\Delta P_{GHG}$  and feedbacks. If  $f$  could eventually approach a  
557 catastrophic value of unity, this ratio reduces to a minimum of 2.

558  
559 In this engineering view, a change in albedo forcing compared with a change in GHGs can be described. The  
560 variation in the energy due to an average albedo change and its re-radiation is

$$561 \quad \Delta P_{\alpha} = \Delta P_{\alpha'} + f_2 \Delta P_{\alpha'} = 1.631 \Delta P_{\alpha'} \quad (\text{A-5})$$

562  
563 The average change in GHGs can be written in terms of  $\Delta f$

$$564 \quad \Delta P_{GHG} = \Delta f P_{GHG} = 1.311\% (f_2 P_{\alpha'}) = 0.827\% P_{\alpha'} \quad (\text{A-6})$$

565  
566 This resulting ratio from Table 1 is

$$567 \quad \frac{\Delta P_{\alpha}}{\Delta P_{GHG}} = \frac{\Delta P_{\alpha'}}{\Delta f} \frac{(1+f_2)}{P_{\alpha'} f_2} = \frac{1.228W/m^2}{0.0131} \frac{1.631}{239.1W/m^2 \cdot 0.631} = 1.01 \quad (\text{A-7})$$

571  
572 Note that this ratio is of course dependent on the 2019 albedo 1.2% change, selected here to obtain unity for  
573 illustrative purposes. The ratio,  $\Delta P_{\alpha'}/\Delta f$ , is an interesting aspect of climate change. In 2019, if we have knowledge of  
574 values, we can compare the dominant aspect of the warming trend. It also provides us with a measure of solar  
575 reversibility

$$576 \quad \Delta P_{\alpha'} \geq \Delta f \frac{P_{\alpha'} f_2}{(1+f_2)} 1.02 \approx 1.21W/m^2 \quad (\text{A-8})$$

578  
579 This ratio is dependent on the change in the albedo compared with a GHG change. It may be helpful in assessing  
580 negative CO2 emissions vs an albedo reduction. Although, it is perhaps not the best way to assess geoengineering

581 estimates. True values of  $\Delta\alpha$  and  $\Delta f$  are not easily obtained in 2019. However, it avoids CO<sub>2</sub> doubling estimates,  
 582 which are also difficult to evaluate. Furthermore, in some instances, a local change in  $\Delta P_\alpha$  can create excess increase  
 583 in GHGs. This has been a concern with cool roofs in the winter which might require additional anthropogenic  
 584 energy. This might be a good way to estimate by Eq. A-8, whether such a change is beneficial by comparison.

585  
 586 It is important to simplify further to provide a more productive approach. In reverse solar geoengineering a global  
 587 warming solution, it is helpful to have simple reliable values. In this view, the 1.6 albedo-GHG factor (which is  
 588 reasonably accurate) is an important engineering number. Another important engineering value is described by a  
 589 Planck-albedo parameter found in Section 5.

## 591 Appendix B: Estimating the Potential for Hotspot irradiance Sensible Heat Storage $H_{T-N}$

592  
 593 A candidate hotspot irradiance sensible heat storage potential  $H_{T-N}$  was described in Section 6. Here we provide a  
 594 preliminary suggested model to clarify and enumerate this factor. It is likely that more rigorous models can be  
 595 developed. Such solutions are outside the scope of this paper.

596  
 597 We consider a ratio for a target (T) area relative to a nominal (N) area defined in Sec. 6. Consider a target area with  
 598 sensible heat storage  $q$  due to a mass  $m$ , having specific heat capacity  $Cp$  experiencing a day-night  $\Delta T$  change in  
 599 time  $\tau$ , then the suggested potential for sensible hotspot heat storage  $H_{T-N}$  has the form

$$601 \quad H_{T-N} = \frac{q_T}{q_N} \times \frac{I_T}{I_N} = \frac{m_T C_{pT} \Delta T_T}{m_N C_{pN} \Delta T_N} \times \frac{I_T}{I_N} \approx \frac{\tau_T C_{pT} \Delta T_T}{\tau_N C_{pN} \Delta T_N} \times \frac{I_T}{I_N} \quad (\text{B-1})$$

602  
 603 Here we provide the option of using temperature change in time  $\tau$  in place of mass. For example, the time to 63%  
 604 change in  $\Delta T$  might be useful (similar to a time constant). We also consider that the irradiance (I) term is needed  
 605 since not all solar absorption energy is stored.

606  
 607 As a numeric example, first consider a 90% irradiance target area (compared to the equator) with nominal mid-  
 608 latitudes (45°) roughly 70%, compared to say the Arctic and Antarctic Circles at 40% [29]. Then the irradiance ratio  
 609 is

$$610 \quad \frac{I\%_T}{I\%_N} = \frac{90\%_T}{70\%_N} = 1.3 \quad (\text{B-2})$$

611  
 612 For the sensible heat numeric portion, consider a rocky area as the target (such as Flaming Mountains). This can be  
 613 compared with a nominal vegetative land area. As a rule of thumb, most rocks have a density of 2.65 g/cm<sup>3</sup>, about  
 614 50% difference compared to a nominal soil area of 1.33 g/cm<sup>3</sup> [32]. The heat capacity of rocks compared with  
 615 vegetated land is 2000 to 830J/Kg/°K [32]. Then  $\Delta T$  is estimated from tables for a day-night cycle [33]. The estimate  
 616 is

$$617 \quad \frac{q_T}{q_N} = \frac{m_T C_{pT} \Delta T_T}{m_N C_{pN} \Delta T_N} = \frac{\rho_T C_{pT} \Delta T_T}{\rho_N C_{pN} \Delta T_N} = \left( \frac{2.65}{1.33} \right)_\rho \left( \frac{2000}{830} \right)_{C_p} \left( \frac{(10^\circ\text{C})}{(6.9^\circ\text{C})} \right) = 2 \times 2.4 \times 1.45 = 6.96 \quad (\text{B-3})$$

618  
 619 Then including irradiance

$$620 \quad H_{T-N} \approx 9 \quad (\text{B-4})$$

## 622 Appendix C: UHI Amplification Factors

623  
 624 An analysis of UHI amplification effects which can be applied to  $H_{T-N}$  was originally provided in Feinberg [5] and  
 625 this work is added here to aid the reader.

### 626 C.1: UHI Area Amplification Factor

627  
 628 To estimate the UHI amplification effects, it is logical to first look at UHI footprint (FP) studies as they provide  
 629 some measurement information. Zhang et al. [34] found the ecological FP of urban land cover extends beyond the  
 630 perimeter of urban areas, and the FP of urban climates on vegetation phenology was 2.4 times the size of the actual  
 631 urban land cover. A more recent study by Zhou et al. [35], looked at day-night cycles using temperature difference  
 632 measurements in China. This study found UHI effect decayed exponentially toward rural areas for the majority of  
 633 the 32 Chinese cities. Their comprehensive study spanned from 2003 to 2012. Zhou et al. describes China as an

634 ideal area to study as it has experienced the most rapid urbanization in the world during the decade evaluated.  
 635 Findings state that the FP of UHI effect, including urban areas, was 2.3 and 3.9 times that of urban size for the day  
 636 and nights, respectively. We note that the average day-night amplification footprint coverage factor is 3.1.

637 The UHI Amplification Factor (AF) is highly complex, making it difficult to assess from first principles as it would  
 638 be some function of

$$639 \quad AF_{UHI \text{ for } 2019} = f\left(\overline{Build}_{Area} \times \overline{Build}_{C_p} \times \overline{R}_{wind} \times \overline{LossE}_{vtr} \times \overline{Hy} \times \overline{S}_{canyon}\right) \quad (C-1)$$

640 were

641  $\overline{Build}_{Area}$  = Average building solar area

642  $\overline{Build}_{C_p}$  = Average building heat capacity

643  $\overline{R}_{wind}$  = Average city wind resistance

644  $\overline{LossE}_{vtr}$  = Average loss of evapotranspiration to natural cooling & loss of wetland

645  $\overline{Hy}$  = Average humidity effect due to hydro-hotspot

646  $\overline{S}_{canyon}$  = Average solar canyon effect

647  
 648 To provide some estimate of this factor, we note that Zhou et al. [35] found the FP physical area (km<sup>2</sup>), correlated  
 649 tightly and positively with actual urban size having a correlation coefficients higher than 79%. This correlation can  
 650 be used to provide an initial estimate of this complex factor. Therefore, as a model assumption, it seems reasonable  
 651 to use area ratios for this estimate.

$$652 \quad AF_{UHI \text{ for } 2019} = \frac{\sum(UHI \text{ Area})_{2019}}{\sum(UHI \text{ Area})_{1950}} \quad (C-2)$$

653 Area estimates have been obtained in the Feinberg [5] yielding the following results for the Schneider et al. [24] and  
 654 the GRUMP [26] extrapolated area results:

$$655 \quad AF_{UHI \text{ for } 2019} = \frac{(Urban \text{ Size})_{2019}}{(Urban \text{ Size})_{1950}} \approx \begin{cases} \left( \frac{[0.188]_{2019}}{[0.059]_{1950}} \right)_{Schneider} = 3.19 \\ \left( \frac{[0.952]_{2019}}{[0.316]_{1950}} \right)_{GRUMP} = 3.0 \end{cases} \quad (C-3)$$

656 Between the two studies, the UHI area amplification factor average is 3.1. Coincidentally, this factor is the same  
 657 observed in the Zhou et al. [35] study for the average footprint. This factor may seem high. However, it is likely  
 658 conservative as other effects would be difficult to assess: increases in global drought due to loss of wet-lands,  
 659 deforestation effects due to urbanization, and drought related fires. It could also be important to factor in changes of  
 660 other impermeable surfaces since 1950, such as highways, parking lots, event centers, and so forth.

661  
 662 The area amplification value of 3.1 is then considered as one of our model assumptions.

663  
 664 **C.2: Alternate Method Using the UHI's Dome Extent**

665  
 666 An alternate approach to check the estimate of Equation C-3, is to look at the UHI's dome extent. Fan et al. [36]  
 667 using an energy balance model to obtain the maximum horizontal extent of a UHI heat dome in numerous urban  
 668 areas found the nighttime extent of 1.5 to 3.5 times the diameter of the city's urban area (2.5 average) and the  
 669 daytime value of 2.0 to 3.3 (2.65 average).

670  
 671 Applying this energy method (instead of the area ratio factor in Eq. C-3), yields a diameter in 2019 compared to that  
 672 of 1950 with an increase of 1.8. This method implies a factor of 2.5 x 1.8=4.5 higher in the night and 2.65 x 1.8=4.8  
 673 in the day in 1950 with an average 4.65. This increase occurs 62.5% of the time according to Fan et al., where their  
 674 steady state occurred about 4 hours after sunrise and 5 hours after sunset yielding an effective UHI amplification  
 675 factor of 2.9. We note this amplification factor is in good agreement with Equation C-3. Fan et al. [36] assessed the  
 676 heat flux over the urban area extent to its neighboring rural area where the air is transported from the urban heat  
 677 dome flow. Therefore the heat dome extends in a similar manner as observed in the footprint studies. If we use the

678 dome concept, we can make an assumption that the actual surface area for the heat flux is increased by the surface  
 679 area of the dome. We actually do not know the true diameter of the dome, but it is larger than the assessment by Fan  
 680 et al. Using the dome extend due to Fan et al. [35] applied to the area of diameter D, the amplification factor should  
 681 be correlated to the ratios of the dome surface areas:  
 682

$$683 \quad AF_{UHI \text{ for } 2019} = \left( \frac{D_{2019}}{D_{1950}} \right)^2 = 2.9^2 = 8.4 \quad (C-4)$$

684 Thus, this equation is a second value for  $H_{T-N}$ , where it is reasonable to use the ratios of the dome's surface area for  
 685 an alternate approach in estimating the effective UHI amplification factor [5]. We will have two values, 3.1 and 8.4  
 686 to work with that provides an upper and lower bounds for effective amplification area.  
 687

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 691

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