

Modeling the Albedo Advantage in Global Warming And an Albedo-Planck Parameter

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Key Words: Re-Radiation Model, Global Warming Modeling, Planck Parameter, Planck-Albedo Parameter

Abstract In this paper, we model global warming using a re-radiation factor and the Planck's parameter to verify consistency. The re-radiation factor is important in quantifying the relative global warming impact of the albedo effect compared to that of greenhouse gases (GHG). Essentially the re-radiation parameter is redefined from the effective emissivity of the planetary system. This parameter found independently in our model has a value of 0.618 (or $\beta=0.887$). The forcing due to the change in the Earth's global albedo compared to GHGs is found to have a 2.6 times larger influence on global warming. In our simple model, we additionally define a handy Planck-Albedo feedback parameter. Using these results, it is concluded that a 1.5% solar geoengineering change in the global albedo could result in -4.8 W/m^2 of forcing. An alternate way to assess the Planck parameter was also found.

1 Introduction

Although global warming is highly complex, often it is helpful to work with a simplified model. We create a model that uses a re-radiation factor which helps to quantify significant differences between changes in the global albedo versus greenhouse gas forcing. It takes into account what normally happens in equilibrium. This is not similar to looking at a comparison of independent feedback parameters $\lambda_{\text{GHG}}/\lambda_{\alpha}$ which provides a different kind of assessment. Here we use a re-radiation parameter obtained mainly in an equilibrium model with appropriate constraints to aid in the comparison; it is then independently found with a unique value of 0.612 (or $\beta=0.887$). This is a redefined variable taken from the effective the emissivity constant of the planetary system. Then the Planck's feedback parameter is used to verify model consistency. This model illustrates a reasonable way to view the Earth's energy budget; it provides a number of useful insights in climatology sensitivity estimates and demonstrates the relative advantage of solar geoengineering solutions over GHG reduction in global warming mitigation [1]. Specifically, a 2.6 larger albedo advantage is found. In working the model, we also find a handy Planck-Albedo parameter that may be useful to climatologists [2] having a convenient value of $1\text{W/m}^2/^{\circ}\text{K}/\Delta\%\text{albedo}$ and this is used to help illustrate the benefits in equilibrium assessments.

2. Data and Method

In order to introduce the re-radiation surface model, it is helpful to initially look at the Planck parameter as it plays a key role in verifying modeling.

2.1 Overview of Planck Feedback Parameter

Estimates on Planck's feedback parameter are varied, typically between $-3.8\text{W/m}^2/^{\circ}\text{K}$ and $-3.21\text{W/m}^2/^{\circ}\text{K}$ with some values as large as $-7.1\text{W/m}^2/^{\circ}\text{K}$ [3]. The IPCC AR4 [4] list a value of $-3.21\text{W/m}^2/^{\circ}\text{K}$. Numerous authors have developed different expressions [3]. A typical estimate starts with

$$F_{\text{TOA}} = (1 - \alpha) S_o / 4 - \sigma(\beta T_s)^4 = (1 - \alpha) S_o / 4 - R_{\text{LWR}} \quad (1)$$

where $S_o=1361\text{W/m}^2$, F_{TOA} is the radiation budget at the top of the atmosphere, R_{LWR} is the outgoing long wave radiation (a function of surface temperature and albedo), σ is the Stefan-Boltzmann constant and β is described in this section below and later will be redefined in terms of a re-radiation parameter. Then the Planck parameter λ_o can be calculated as

$$\lambda_o = \partial F_{\text{TOA}} / \partial T_s = -\partial R_{\text{LWR}} / \partial T_s \quad (2)$$

This result is

$$\lambda_o = -4\beta^4 \sigma T_s^3 = -4\beta \sigma T_{\text{toa}}^3 \quad (3)$$

where β varies in the literature from 0.876 to 0.887 (averaging=0.8815) and $T_s=288^{\circ}\text{K}$ [4]. This yields $-3.37\text{W/m}^2/^{\circ}\text{K} < \lambda_o < -3.21\text{W/m}^2/^{\circ}\text{K}$.

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62 However, from Eq. 3, β is often taken as the ratio

$$63 \quad 64 \quad \beta = T_{TOA} / T_s = 255^\circ K / 288^\circ K = 0.8854 \text{ and } \beta^4 = 0.615 \quad (4)$$

65
66 A common assessment uses $T_{TOA}=255^\circ K$, so that $\lambda_o = -3.33 W/m^2/^\circ K$. Another expression developed by Schlesinger
67 [5] is dependent on the albedo and surface temperature as

$$68 \quad 69 \quad \lambda_o = S_o (1 - \alpha) / T_s \quad (5)$$

70
71 When $S_o=1361$, $0.294118 < \alpha < 0.3$, and $T_s=288^\circ K$ then $-3.308 W/m^2/^\circ K > \lambda_o > -3.3358 W/m^2/^\circ K$, respectively.

72 **2.2 Estimating Planck's Parameter with an Albedo Method**

73
74 Consider a global albedo change corresponding to $1^\circ K$ rise from solar absorption. Since we are only concerned with
75 an albedo change

$$76 \quad 77 \quad 78 \quad F_{TOA} = 0 = (1 - \alpha) E_o - \sigma (T_s)^4 \quad (6)$$

79 where $E_o = S_o/4$. Then a $1^\circ K$ change is

$$80 \quad 81 \quad \Delta T_s = T_2 - T_1 = \left(\frac{E_o}{\sigma} (1 - \alpha_2) \right)^{1/4} - \left(\frac{E_o}{\sigma} (1 - \alpha_1) \right)^{1/4} = 1^\circ K \quad (7)$$

82
83 Here we will use the AR5 albedo starting value of 0.294118 [6]. We find that the corresponding albedo change is
84 0.28299 when $E_o=340 W/m^2$. This corresponds to an absorption of

$$85 \quad 86 \quad \Delta E_o = E_o \{ (1 - \alpha_2) - (1 - \alpha_1) \} = E_o (\alpha_1 - \alpha_2) = 3.784 W / m^2 \quad (8)$$

87
88 Since this is for a $1^\circ K$ rise, then it can also be written as

$$89 \quad 90 \quad \lambda_{1K} = 3.784 W/m^2/^\circ K \quad (9)$$

91
92 We note this is related to the surface value, then

$$93 \quad \lambda_{1K} = -4\sigma T_s^3 \quad (10)$$

94 By comparison to above we have

$$95 \quad \lambda_o = \lambda_{1K} \beta = -3.784 W/m^2/^\circ K = -3.349 W/m^2/^\circ K \quad (11)$$

96
97 This is very close to the $-3.33 W/m^2/^\circ K$ value obtained in the traditional manner.

98 **2.3 Top of the Atmosphere and Beta**

99
100 From Eq. 1

$$101 \quad 102 \quad R_{LWR} = \sigma (\beta T_s)^4 = \sigma (T_{TOA})^4 \quad (13)$$

103
104 giving

$$105 \quad \beta^4 R_{TOA, T_s} = R_{TOA, T_{TOA}} \quad (14)$$

106
107 We will need this expression later when showing model consistency with the Planck feedback parameter.

108 **2.4 Re-radiation GHG GW Model**

109
110 In this model we define

$$111 \quad 112 \quad P_{Total} = \sigma T_s^4 \text{ and } P_\alpha = \sigma T_\alpha^4 \quad (15)$$

113
114 We consider a time when there is no feedback issues. Then by conservation of energy, the equivalent power re-
115 radiated from GHGs in this model is

$$116 \quad 117 \quad 118 \quad P_{GHG} = P_T - P_\alpha = \sigma T_s^4 - \sigma T_\alpha^4 \quad (16)$$

119

120 Since typically, $T_\alpha \approx 255^\circ\text{K}$ and $T_s \approx 288^\circ\text{K}$, then we note in keeping with the definition of Beta (see Eq. 4) for the
 121 moment, that $\beta \approx T_\alpha/T_s$. This allows us to write
 122

$$123 \quad P_{GHG} = \sigma T_s^4 - \sigma T_\alpha^4 = \frac{\sigma T_\alpha^4}{\beta^4} - \sigma T_\alpha^4 = \sigma T_\alpha^4 \left(\frac{1}{\beta^4} - 1 \right) \quad (17)$$

124 We note that when $\beta^4=1$, there are no GHGs as required by definition of β . We now define a re-radiation parameter
 125 $f = \beta^4$. We know that some fraction of the blackbody radiation is re-radiated by the GHGs so f is a re-radiation
 126 parameter. That is, the energy, P_{GHG} , must be some fraction P_α so that
 127

$$128 \quad P_{GHG} = f P_\alpha = f \sigma T_\alpha^4 \quad (18)$$

129 However, in order for this to be true it requires
 130

$$131 \quad P_{GHG} = \sigma T_\alpha^4 \left(\frac{1}{f} - 1 \right) = f \sigma T_\alpha^4 \quad (19)$$

132 This leads us to solutions of the quadratic equation
 133

$$134 \quad f^2 + f - 1 = 0 \text{ yielding } f = 0.618034 = \beta^4, \quad \beta = (0.618)^{1/4} = 0.88664 \quad (20)$$

135 This is very close to the value estimated for β and was obtained though energy balance in the planetary system
 136 providing a completely independent assessment without any approximations. In Section 2.6, we double check in
 137 another way by balancing energy in and out.
 138

139 2.5 Re-radiation Model Applied to Two Different Time Periods

140 Global warming can be modeled by looking at two different time periods. We can model the radiation for 1950 as
 141 due to blackbody radiation with the addition of GHG re-radiation where in this time period

- 142 • we will assume no feedback issues causing a warming trend so that

$$143 \quad P_{Total, 1950} = P_\alpha + P_{GHG} = P_\alpha + f_1 P_\alpha \quad (21)$$

144 where $P_\alpha = S_o \{0.25x(1 - Albedo)\}$ and $S_o = 1361 \text{W/m}^2$. The equilibrium model is constrained by energy balance
 145 discussed in Section 2.4 and 2.6. In 2019 due to global warming trends, this model is more complex and harder to
 146 separate out terms. However, it can still be done looking at a snapshot point in time using equilibrium theory, so

$$147 \quad P_{Total 2019} = P_{\alpha'} + P_{GHG' + Feedback} = P_{\alpha'} + f_2 P_{\alpha'} \quad (22)$$

148 Here $P_{GHG' + Feedback}$ includes GHGs and its increase comprising also of water-vapor increase, lapse rate feedback and
 149 other effects such as an increase in snow-ice albedo changes that are hard to separate out. That is, some of this
 150 feedback is related to GHG increases and some is related to albedo change. $P_{\alpha'}$ represents any albedo change due to
 151 UHI absorption increases, cloud absorption change, ice and snow melting and so forth that can be discerned. We
 152 note that f , a measure of the emissivity, is **not** constant but must change since the amount of GHGs change.
 153

154 However, the re-radiation still must connect the absorption to re-radiation. We have used a linear f parameter that
 155 indicates the fraction of P_α power that must be re-radiated back to obtain the observed temperature. To be clear, f is
 156 just a fractional parameter related to the emissivity. In 1950 it is some function of the GHGs (with no feedbacks). In
 157 2019 it is more complex. The model is also constrained relative to f_1 as described in Section 2.6. However, it is
 158 primarily related to GHGs re-radiation since $P_{GHG} \approx P_{GHG' + Feedback}$.

159 2.6 Balancing P_{out} and P_{in}

160 Although Eq. 15 is reasonably simple, it turns out that f_1 has a uniquely defined value obtained when balancing the
 161 energy.
 162

177 2.6.1 Balancing P_{out} and P_{in} in 1950

178
179 In order to balance the energy in with the energy out in 1950 with no global warming imbalance we can still start
180 with Eq. 21. In equilibrium the radiation that leaves must balance what comes in P_α so that
181

$$182 \quad \begin{aligned} Energy_{out} &= (1-f_1)P_\alpha + (1-f_1)P = (1-f_1)P_\alpha + (1-f_1)\{P_\alpha + f_1P_\alpha\} \\ &= (1-f_1)\{2P_\alpha + f_1P_\alpha\} = 2P_\alpha - f_1P_\alpha - f_1^2P_\alpha = Energy_{in} = P_\alpha \end{aligned} \quad (23)$$

183
184 In 1950 the value of f solves the quadratic equation
185

$$186 \quad f_1^2 + f_1 - 1 = 0 \quad \text{yielding } f_1 = 0.618 \quad (24)$$

187 Interestingly, this also says that
188

$$189 \quad P_\alpha = f_1 P_{Total_1950} \quad \text{or } P_\alpha = f_1(P_\alpha + f_1P_\alpha) \quad \text{or } 1 = f_1(1 + f_1) \quad (25)$$

190
191 The RHS of Eq. 25 is Eq. 24 and Eq. 20. This is why f_1 is unique. It is the fractional amount of total radiation that is
192 in equilibrium. As a final check, results will show in Section 3 and Table 1, that the value f_1 provides reasonable
193 results.
194

195 2.6.2 Warming Imbalance in 2019

196
197 The re-radiation parameters f_1 and f_2 are connected and from Eq. 21 and 22 we have
198

$$199 \quad f_2 = f_1 + \left(\frac{P_{2019}}{P_{\alpha'}} - \frac{P_{1950}}{P_\alpha} \right) = f_1 + \Delta f \quad (26)$$

200 In this way f_2 is a function of $f_1=0.618$ and the differences in the global warming residuals that is defined in Eq. 26
201 as Δf .
202

203 3.0 Results and Discussion

204
205 Since the re-radiation parameter $f_1=0.618$, in order to obtain $T_{1950}=13.89^\circ\text{C}$ (287.038°K), the only adjustable
206 parameter in our simple model is the Earth's albedo. This value requires an albedo value of 0.3008 (see Table 1) to
207 obtain the correct value T_{1950} . This is a reasonable and similar to values cited in the literature [11].
208

209 In 2019, the average temperature of the Earth is $T_{2019}=14.84^\circ\text{C}$ (287.99°K). Here we are not sure of the albedo since
210 it likely changed due to UHI increase, snow and ice melting and cloud coverage changes. The IPCC value in AR5
211 [6] is 0.294118. However, this would represent a 3% change since 1950 which may be an overestimation. In our
212 assessment, we will assume a 1% change. Then the f_2 parameter is adjusted to 0.6324 in order to obtain T_{2019} .
213 Results are provided in the Table 1. The results yields $P_{Total_1950}=384.918 \text{ W/m}^2$ and $P_{Total_2019}=390.024 \text{ W/m}^2$. We
214 find that

$$215 \quad \Delta P_{Total} = P_{2019} - P_{1950} = 5.097 \text{ W/m}^2 \quad (27)$$

216 and

$$217 \quad \Delta T_{Total} = T_{2019} - T_{1950} = 0.95^\circ\text{C} \quad (28)$$

218 which is the observed surface temperature increase since 1950.
219
220

Table 1 Model results

Year	T($^\circ\text{K}$)	T_α ($^\circ\text{K}$)	f_1, f_2	α, α'	$P_\alpha, P_{\alpha'}$ (W/m^2)	$P_{\text{GHG}}(\text{W/m}^2)$ $P_{\text{GHG}'+\text{feedback}}$	P_{Total} (W/m^2)
2019	287.991	254.78	0.63253	29.779	238.927	151.128	390.055
1950	287.041	254.51	0.6180	30.08	237.903	147.032	384.935
$\Delta 2020-1950$	0.95	0.27	1.45%	-0.3 (1%)	1.024	4.096	5.121

221
222 The table below summarizes model results for the specified albedos and setting the model to the observed Earth's
223 surface temperatures.
224

225 To show model consistency, the forcing change 5.121 W/m^2 resulting in a 0.95 $^\circ\text{K}$ rise, should agree with what is
226 expected from Planck's feedback parameter. From Eq. 14 it is evident that
227

$$\beta^4 \Delta R_{TOA} = 5.121 \times \beta^4 = 3.1 \text{ W/m}^2 \quad (29)$$

228
229
230 This illustrates the consistency of the simple re-radiation model. Then Planck's feedback parameter ($3.3 \text{ W/m}^2 / ^\circ\text{K}$)
231 temperature rise is in reasonable agreement with what is observed by equilibrium modeling

$$3.165 \text{ W/m}^2 \times (1/3.3) ^\circ\text{K/W/m}^2 = 0.959 ^\circ\text{K} \quad (30)$$

235 3.1 Why the Re-radiation Parameter is Significant

236
237 In Table 1, the measure of $\Delta f = 1.45\%$ fractional increase is due to re-radiation change. This is significant. From Eq.
238 21, 22 and 26 we can illustrate this key characteristic of the climate change

$$\Delta f = \left(\frac{P_{2019}}{P_{\alpha'}} - \frac{P_{1950}}{P_{\alpha}} \right) = \left(\frac{P_{GHG'+F}}{P_{\alpha'}} - \frac{P_{GHG}}{P_{\alpha}} \right) \approx \left(\frac{P_{GHG'+F} - P_{GHG}}{P_{\alpha}} \right) \quad (31)$$

241
242 Therefore f is an estimate of climate re-radiation and Δf an estimate of climate emissivity change. It is a measure of
243 GHG increase and the feedback relative to the initial radiation, and is generally helpful in looking at how our
244 climate is working. Furthermore, we can deduce an albedo advantage.

246 3.2 The Albedo Advantage

247
248 We can look at an important ratio, the power created by the albedo effect compared to GHGs in 1950. The initial
249 radiation is P_{α} which heats the Earth to $254.51 ^\circ\text{K}$ then according to Eq. 21 and Table 1, the P_{GHG} energy originates
250 from a fraction of this original heating due to re-radiation as $f P_{\alpha}$

$$\frac{P_{\alpha} + P_{GHG}}{P_{GHG}} = \frac{P_{\alpha} + f P_{\alpha}}{P_{GHG}} = \frac{P_{\alpha} + f P_{\alpha}}{f P_{\alpha}} = \frac{1 + f_1}{f_1} = \frac{1.62}{0.62} = 2.62 \quad (32)$$

253
254 In general, this also means that albedo change has a higher impact factor in climate forcing, 2.6 times larger than
255 ΔP_{GHG} as well, that is a change, ΔP_{α} compared with a change in ΔP_{GHG} would yield the same impact factor
256 $d(P_{\alpha} + P_{GHG}) = 2.62 d(P_{GHG})$ or assuming $\Delta f \ll 1$

$$\frac{\Delta P_{\alpha} + \Delta P_{GHG}}{\Delta P_{GHG}} \approx \frac{\Delta P_{\alpha} + f \Delta P_{\alpha}}{f \Delta P_{\alpha}} \approx \frac{1 + f_1}{f_1} = \frac{1.62}{0.62} = 2.62 \quad (33)$$

259
260 This is a key reason that UHIs, cloud coverage, snow and ice melting, can create significant climate effects.
261 Appendix A puts this important impact factor in layman's terms. We see this is a different kind of comparison then
262 $\lambda_{GHG}/\lambda_{\alpha}$. It uses a re-radiation emissivity parameter obtained mainly from the equilibrium model.

263
264 In this view, an albedo solution is advantageous having significant potential for reversing global warming or
265 ignoring it, as in UHIs and roads, likely can create serious issues. Therefore, trying to control global warming by
266 reducing GHGs is important. However, certainly an albedo approach is more advantageous. It reduces both initial
267 absorption and its potential for its re-radiation. Its impact rating can be taken as 162% compared to re-radiation f
268 with a 62% impact by comparison according to Eq. 32 and 33, yielding a 2.6 times higher advantage. It is important
269 to realize that because the albedo solution can highly impact GW and reverse trends, it is also vital in preventing a
270 tipping point from occurring.

272 3.3 Planck-Albedo Feedback Parameter

273
274 The albedo and ΔP_{α} change in Table 1, is: $\% \Delta \alpha = 1\%$ and 1.024 W/m^2 , respectively. We note this defines a unique
275 Planck-Albedo parameter $\lambda_{\% \Delta \alpha} = \Delta P_{\alpha} / \% \Delta \text{albedo}$. To illustrate from Table 1

$$\lambda_{\% \Delta \alpha} = 1.024 \text{ W/m}^2 / \Delta \% \text{albedo} = 1.024 / 1\% \quad (34)$$

278
279 This parameter can also be expressed per degree (noting the $0.95 ^\circ\text{K}$ change in Table 1)

$$\lambda_{\% \Delta \alpha \Delta T} \approx 1 \text{ W / m}^2 / \Delta \% \text{albedo} / ^\circ\text{K} \quad (35)$$

282
283 The parameter was first noted in Feinberg 2020 [2] but is featured here as a modeling tool. We term it the Planck-
284 Albedo parameter, since it relates to blackbody (P_{α}) absorption. A simple numeric example is given in the

285 conclusion to illustrate how it can provide helpful estimates. This interesting parameter arises from the basic
 286 assessment of the two equilibrium time periods

$$287 \quad \lambda_{\omega_{\Delta\alpha}} = \frac{(\Delta E_o)_{\alpha}}{\alpha_1 - \alpha_2} \frac{100}{\alpha_1} = \frac{E_o (\alpha_1 - \alpha_2)}{\alpha_1 - \alpha_2} \frac{100}{\alpha_1} = E_o \alpha_1 / 100 \approx 1W / m^2 / \% \Delta albedo \quad (36)$$

288 where $E_o=340 \text{ W/m}^2$ and when α_1 is 29.4118%, the value of $1.000\text{W/m}^2/\Delta\%albedo$ is obtained. We note the value
 289 29.4118% (100/340) is given in AR5 [6]. The parameter's relationship to λ_{α} is

$$291 \quad \lambda_{\alpha} = \lambda_{\omega_{\Delta\alpha\Delta T}} x \% \Delta \alpha \quad (37)$$

292 and the feedback parameter including f re-radiation is in 2019
 293

$$295 \quad \lambda_{\alpha}^{\dagger} = \lambda_{\omega_{\Delta\alpha\Delta T}} x \% \Delta \alpha x 1.618 \quad (38)$$

297 4.0 Conclusion

299 In this paper we provided a simple re-radiation global warming model. The model shows consistency with the
 300 Planck parameter. We noted that the re-radiation parameter increased by about 1.45% due to global warming from
 301 1950 to 2019, illustrating the warming from a different perspective. From the model, the albedo effect was
 302 quantified having an impact rating of 162% compared to GHGs with 62%. The albedo effect then yields a 2.6 times
 303 higher advantage upon comparison. These results strongly support moving forward with solar geoengineering
 304 solutions [2, 7-9].

306 We also found a handy parameter that we termed the Planck-Albedo parameter which is about
 307 $\lambda_{\omega_{\Delta\alpha\Delta T}} \approx 1W / m^2 / \Delta\%albedo / ^{\circ}K$. This can be helpful in quickly estimating the effect of an albedo change on global
 308 warming and in assessing λ_{α} . For example, Feinberg 2020 [1] suggested a goal of 1.5% geoengineering albedo
 309 change. Using this parameter, an impact of 1.5 Watts/m² warming reduction should result. Given a 1.62 reemission
 310 factor (Eq. 32), this is 2.4W/m² improvement. With a reduction in water-vapor feedback, often estimated by a factor
 311 of 2 [10], provides an overall resulting effect that could be as high as 4.8W/m². Feasibility is discussed in more
 312 detail in Feinberg's 2020 paper [1] and other solutions have been proposed [6-9].

314 Appendix A: Quantifying the Albedo Advantage in Layman's Terms

316 It may be helpful for the reader to have a layman's view of the 2.62 factor. Consider the Earth with a roof. The roof
 317 represents the GHGs over the Earth and only allows 40% of any energy to leave with the rest returning to Earth.
 318 Sunlight comes in and some is absorbed and heats the Earth's floor to 255°K (-2.3°F very cold). Let's say it takes
 319 100 units of energy. The heat rises but only 40 units of energy can leave from the roof, so 60 units comes back and
 320 warms the Earth's floor some more to 288°K (57°F average temp of Earth). On average the Earth's floor is warmed
 321 a total of 160 units. The Sun keeps warming the Earth's floor at 100 units on average and the roof keeps sending
 322 back 60. So the roof is responsible for 60 units on average of energy and the Earth's floor is warmed up to 160 units
 323 on average. We can write this as

$$325 \quad \text{Energy units: } 160=100+60=100+100 \times 0.6$$

327 We see the 100 units is in two places in the equation due to the floor and roof, while the 60 is only in one place. That
 328 is, without the floor absorption first, the roof cannot keep the Earth warm. Therefore, the heat coming from the
 329 Earth's floor results in 160 units and the roof is only 60 units by comparison. The impact factor is

- 331 • 160/60=2.66, that is the heat from the Earth's floor has this much larger impact.

333 Alternately, for every unit of energy given off by the Earth's floor after absorption, it is equivalent to causing 1.6
 334 units of heating while the roof (GHG) is only responsible for 0.6.

336 How much heat leaves in equilibrium? There was the initial 40 leaving of the 100 units of energy absorbed and
 337 radiated. As well the Earth's floor received a total of 160 units but the roof only let 40% leave that is another 64
 338 (=0.4 x 160) units of energy leaving. The total leaving is 104 units in equilibrium so roughly 100 units comes in and
 339 almost same goes out.

341 This can be refined to 61.8% (Eq. 20). Then 100 units is absorbed and radiated, then 38.2 units initially leave, and
 342 61.8 units is radiated so the Earth's floor is heated to 161.8 units of energy. From this 0.382 x 161.8 leaves=61.8
 343 units or energy. The total is 61.8+38.2=100 units of energy leaves and another 100 units comes and equilibrium is
 344 established. Any difference causes global warming.

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